Permian–Triassic riftogenic processes

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This contribution is aimed to:

- 1) The Permian and Triassic meta-magmatics investigated in the Inner Western Carpathians and Austro-Alpine units (geology, petrology, geochemistry, geochronology);
- 2) The geodynamic background of the continuous Permian–Triassic tectonomagmatic events

Methods:

field (structural), laboratory (petrotectonic, EPMA, micro-Raman, LA-ICP-MS, ICP-MS, EPMA Mnz dating, SIMS and nano-SIMS isotopic U/Pb, and LA-MC-ICP-MS dating of Zrn, Rt, Nd-Sr-Pb isotope systematics

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Permian-Triassic Plum generation Zone (PGZ) in red circle



Fig. 11.2 (a) CEED mantle frame reconstruction of the estimated eruption Central Atlantic Magmatic Province (CAMP) together with 38 Triassic kin contours in the lower mantle (Lekic et al., 2012). Areas 5, 3, and 1 define mantle. The 1% slow SMEAN contour is shown for comparison (see Cha locations erupted near or over the Tuzo Plume Generation Zone (PGZ) w (231–215 Ma) appear to be related to the Perm Anomaly. Outline Earth ge (c) 210 Ma (Late Norian). A, Annamia; F, Falkland; K, Kazakh terranes; SC, South China; T, Tarim; TH, Tethyan Himalaya. CEED palaeomagnetic







(Putiš et al., 2018, Lithos)

websterite GS-ST

ε_{Hf}(t) zircon values of 16-8 (average 11 from 18 spots) suggest a depleted mantle source of these sub-crustal cumulate-type melts

Steinbach Clinopyroxenite dyke in harzburgite: Permian magmatic age 252±2.4 Ma Zircon U-Pb SIMS (UAA HP Sieggraben str. c.) Putiš et al. (2018, Lithos)

Upper A-Alpine hanging wall HP Sieggraben structural complex

P–T calculated for D1 stage peak pressure conditions in clinopyroxenite by Perple_X

Reconstructed P–T path of pyroxenites in harzburgite block (Putiš et al., 2018, Lithos)

Infratatric sedimentary cover (Považský Inovec Mts.)

280

U/Pb SIMS Permian 267±2 Ma calc-alk. Meta-rhyolite intercalation within arkosic sandy shales (Putiš et al., 2016)

rhyolite

260

The Gemeric Superunit Permian granites

granit Humel

The Gemeric Permian granites (magmatic ages in red)

record a

Triassic overheating by the green Mnz ages up to the Carnian-Norian,

which may be related to the opening of the Meliata Basin during Advanced continental (~250-240 Ma) and ncipient oceanic (from~240 Ma)

rifting stages

Radvanec, Németh, Putiš et al. (2009)

Basaltic tuff

Radiolarite JAK-2/2

Cherty clay-carb. shale JAK-2/1

Carbonate JAK-2/3

<mark>5 cm</mark>

Peperitic breccia horizon

Detrital Zircon in Ladinian cherts (the Meliatic oceanic Jaklovce Succession), with ca. 0.5 m thick N-MORB layers

Harzburgite dark core (s. DO-2d)

OI2(±Px2)

Opx1 (with incl. of Ol1, Spl1)

01

Cpx2

<u>Ultramafics</u> from the Dobšiná serpentinite mélange (Meliaticum)

Serpentinite
Prv-bearing zone

Prv-free zone harzburgite

Weakly depleted abyssal Spl Harzburgite

Putiš et al. (2016)

<u>Jaklovce Unit</u>: JAK-30, JAK-204B, JAK-6/19; <u>Meliata Unit</u>: JAK-204C, MEL-11/19 (olistolith of Bôrka Unit); Sillica Nappe: SA-6.

Putiš et al. (2019) Nemec et al. (2020)

Discussion

 Pyroxenites and websterites, related to sub-crustal cumulate-type melt derivates (less evolved for websterites and highly evolved for clino- and orthopyroxenites) in the AA Unit provide an evidence about the Late Permian/Early Triassic (~252 Ma) mantle melting.
 Permian–early Middle Triassic (Anisian) magmatism (~280–245 Ma) from the continental crust AA and IWC units belong to the same tectono-magmatic event which culminated with the Neotethyan oceanic Meliata Basin opening in the Middle Triassic (Ladinian). <u>Preliminary Conclusion</u>

> The major Permian–Triassic magmatism exceeded the inferred active Paleotethyan margin and continued W-ward along the Paleotethyan suture zone. This unstable zone suffered from a longterm (from ca. 300 to 200 Ma) hot lithosphere, where the Late Variscan orogen post-collisional collapse continued by thermal erosion of the crustal and mantle lithosphere due to the asthenosphere upwelling and thermal perturbation through the extensionally thinned continental crust. A support from the mantle plum generated from Paleozoic subductions can be considered.